

# Fjernmåling. Jordobservasjoner - Envisat

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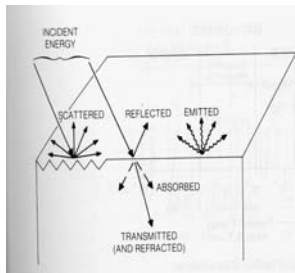
## Fjernmåling av atmosfæren

Kan foregå fra bakken eller rommet (ballong, fly, satellitt, romferge)

- 3 prinsipper
- Målemetoder (delvis dekket av Stette i kap. 7)
- Anvendelser (miljøovervåkning, meteorologi)

## 3 teknikker

- Passiv måling: benytter sol lys som strålingskilde, kamera eller spektrofotometer som detektor.
- Passiv måling: benytter varmestråling fra jorda som strålingskilde, kamera eller spektrofotometer som detektor av re-emittert stråling fra atmosfæren
- Aktiv fjern måling (bruker strålingskilde (laser, radar, mikrobølgegenerator) og detekterer refleksjoner og spredt lys.

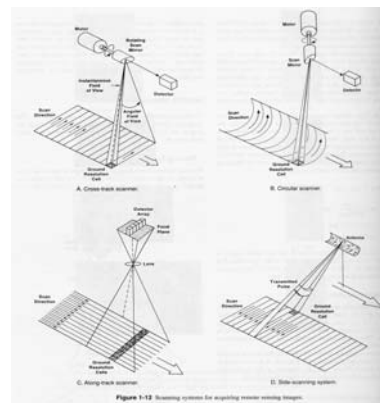


## Detektorteknologi

- Solar radiation detectors
  - detector
    - CCD cameras (diodes)
    - photomultipliers (Spectroradiometers)
  - filters
    - bandwidth narrow <10 nm
    - broad band 20-100 nm
  - field of view and direction (nadir)
  - different scanning modes (limb)
  - solar radiation (0.2 – 3 m)

- Radar
  - microwave transmitter and receiver
  - wavelength 100 – 1 cm
  - active (transmitter on the satellite)
- Lidar
  - light transmitter (laser) and receiver
  - active sensing, only
- Thermal detector
  - ”temperatur sensors” from –20 Celsius to several hundred degree Celsius
  - wavelength 1000 – 7 m

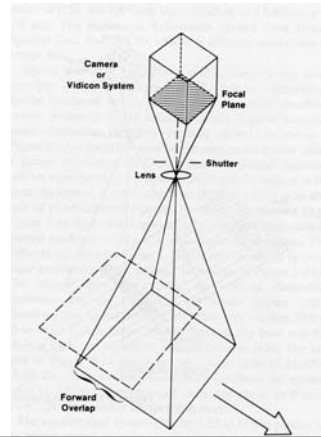
## Tracking techniques



# Billedbehandling

- Preprocessing
  - radiometric (wavelength bands, brightness, remove "background" signal)
  - image matching (geometric correction by resampling)
  - image classification (meaning of each pixel different correlation procedures – neural network)
- Validation
  - field data (from ground)
  - other satellittes

# Overlapping



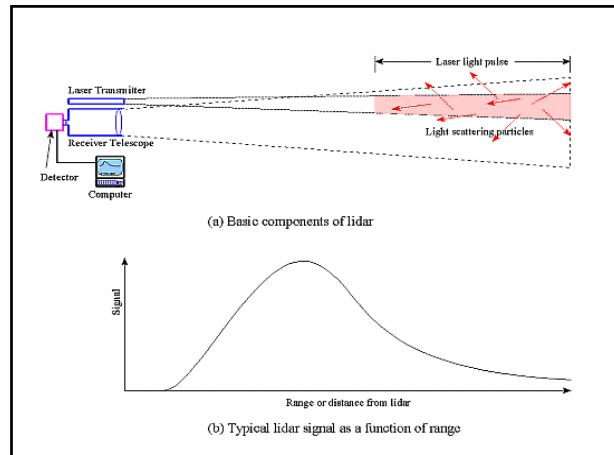
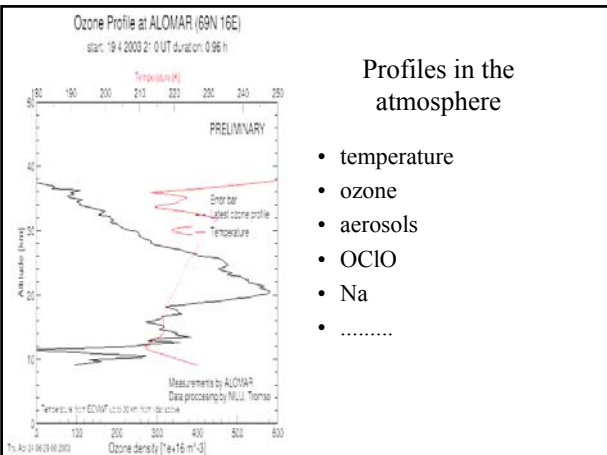
# Aktiv fjernmåling

LIDAR  
Fra bakken og fra rommet

# Lidar Light Detection And Ranging (similar to RADAR)

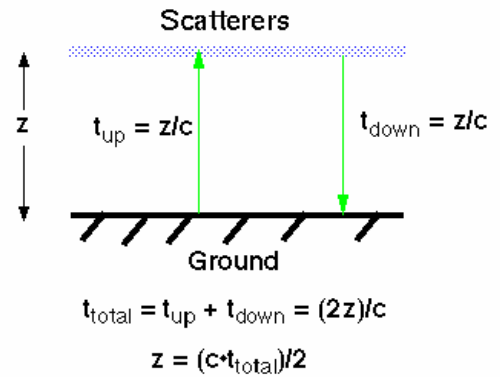


Bilde fra Alomar ved Andøya raketstasjon, Lidar



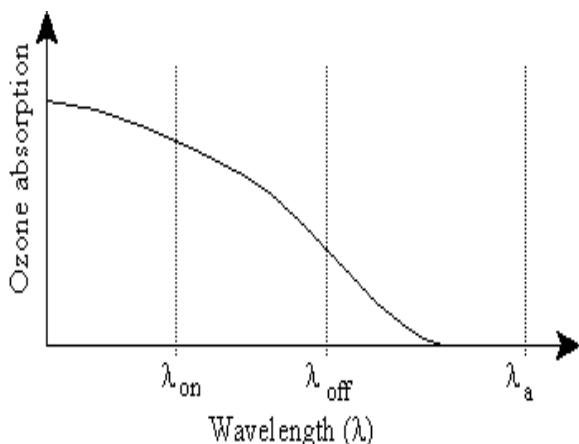
A lidar transmits short pulses of laser light into the atmosphere. The laser beam loses light to scattering as it travels. At each range, some of the light is backscattered into a detector. [Fig. 1a]

Because the light takes longer to return from the more distant ranges, the time delay of the return pulses can be converted to the corresponding distance between the atmospheric scatterer and the lidar. The end result is a profile of atmospheric scattering versus distance. [Fig. 1b] Analysis of this signal can yield information about the distribution of aerosols in the atmosphere. The amount of backscatter indicates the density of the scatterers.

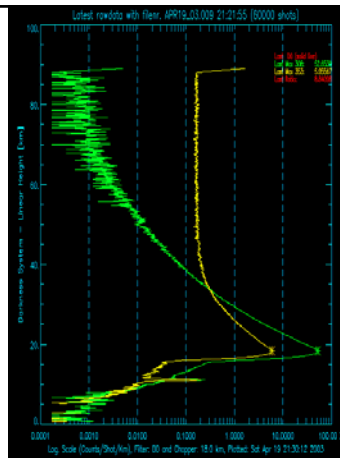


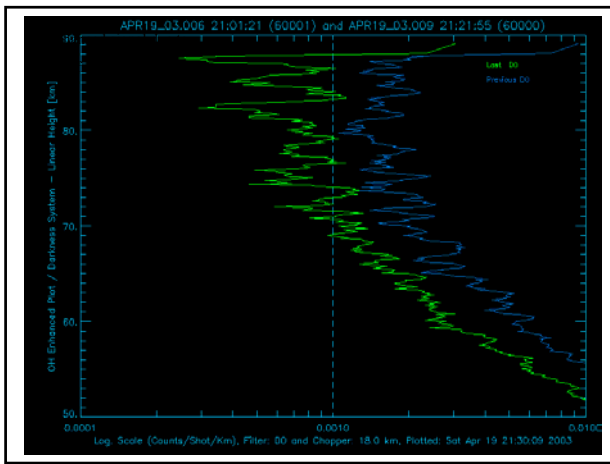
- A frequency shift in the light because of the Doppler effect permits measurement of wind speeds.
- By detecting the amount of depolarization, one can discriminate between liquid droplets and nonspherical ice particles.
- Differential Absorption Lidar (**DIAL**) uses absorption, as evidenced by reduced backscatter from greater distances, to measure the concentration of atmospheric gases.
- A Raman lidar detects particular atmospheric components (such as water vapor) by measuring the wavelength-shifted return from selected molecules.

- DIAL measurements of ozone require less precise laser frequency because ozone has a broad (~200 nm) absorption band instead of narrow lines like water vapor. This requires that the on- and off- wavelengths be chosen with sufficient wavelength separation to ensure a significant difference in their absorption.
- However, uncertainty in the changes of aerosol scattering and patterning between the wavelengths can introduce error. The difference in aerosol scattering at  $\lambda_{on}$  and  $\lambda_{off}$  can be extrapolated from a third measurement taken at a longer wavelength,  $\lambda_a$ .
- Therefore, DIAL measurements with three wavelengths can be used to determine ozone concentration profiles to good accuracy.



Reference:  
<http://alomar.rocketrange.no/ozon-latest-rawdata.html>  
 Profiles of 308 nm and 353 nm laser light  
 Algorithms give ozone profiles





## Passiv fjernmåling

Reflektert solstråling fra jorden og atmosfæren målt med instrumenter på satellitt

**Satellitt: miljø overvåkning**

1960 - Første værssatellitt kun til meteorologi (Tiros1)

1970 - Første miljøovervåkningsinstrument (ozon, BUV) Nimbus 4


1972 - Landsat (jord observasjoner) NASA

1984 - Første rene klimaovervåknings satellitt. ERBS (Earth Radiation Budget Satellite, ERS 1 og 2)


2002 - Envisat

## Bane for miljøovervåknings-satellitter

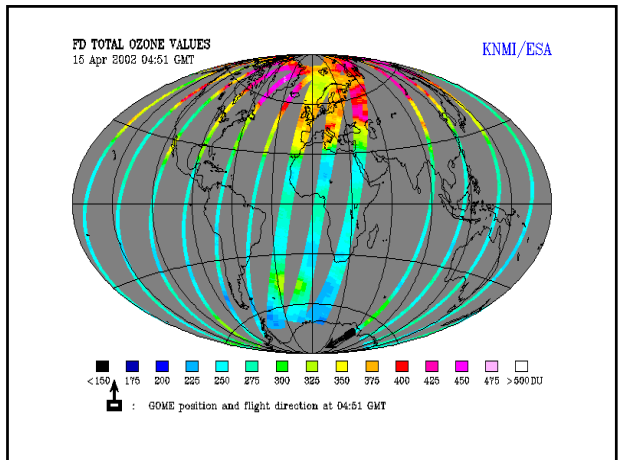
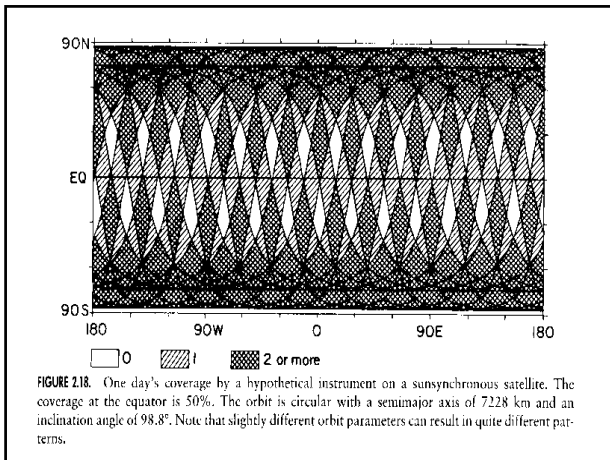
Satellitt banen sett fra et punkt som roterer med jorden

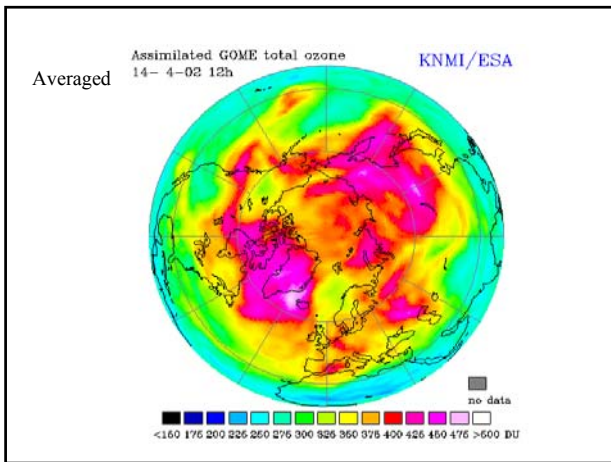


Tidsforløp for målinger



- Meteo - satellittene - GEO
- Klima - satellittene - POLAR
- Høy "inclination angle" - 96°-99° eller Low Earth orbit
- Dekker nesten hele kloden på et døgn.
- Deler av ekvator blir ikke dekket.
- Stort sett for man en observasjon pr. døgn pr. sted

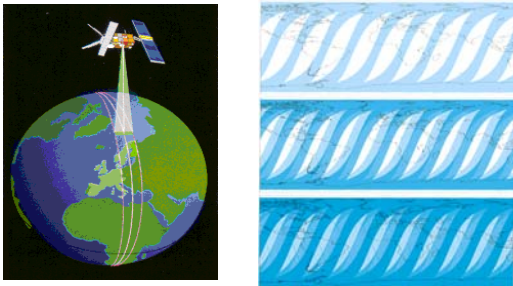




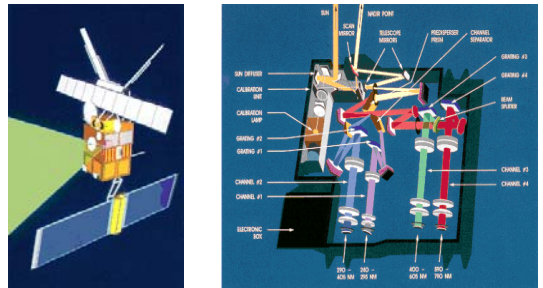
## ERS 2 - ESA environmental satellite

- ERS-2 flies in a sun-synchronous polar orbit with an inclination of  $98^\circ$ , a mean altitude of 780 km, and a local time of 10.30 h of the equator crossing at the descending node. This results in an orbital period of approximately 100 minutes and a speed of the subsatellite point of 7 km/s.
- The instrument instantaneous field of views has been chosen to be  $40 \times 2$  km, or  $2.8^\circ \times 0.14^\circ$ .
- By movement of the scan mirror, this instantaneous field of view can be scanned across track. With a  $\pm 31^\circ$  scan, a global coverage can be obtained within 3 days, except for a gap of approximately  $4^\circ$  around both poles. To overcome this limitation a special pole view mode has been included.

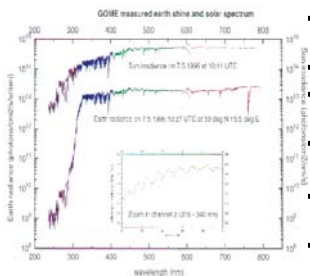
## ERS 2 - covers the earth in 3 days



## ERS2 - Gome UV and visible spectroradiometer

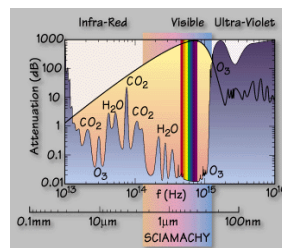


## Instrumentering - optiske instrumenter



- Spektral instrumenter likt med bakkeinstrumenter
- Inngangsoptikk
- Monokromator (med gitter)
- Detektor
- Måler spekteret fra solen (øverste kurve)
- Måler spekter tilbakereflektert fra atmosfæren (nederste spekter)

## Fjernmålingsalgoritme: fra strålingsmåling til mengde gass i atmosfæren.



- Stråling i atmosfæren
  - absorpsjon
  - spredning
  - refleksjon
- Algoritmer som omdanner solstrålingsmålingene til mengde gass.
- Mest benyttet: måling av absorpsjon ved to nærliggende bølgelengder.
- Forskjell i intensitet angir mengde absorberende stoff

## Beer Lambert's law; Absorption

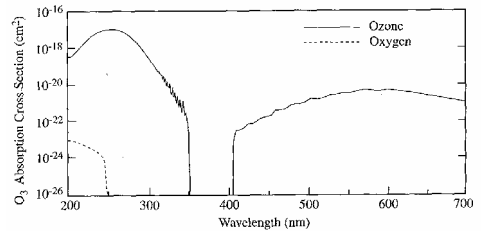
The *optical depth* of a medium at a given distance  $z$  from 0 is defined by:

$$\tau = \int_0^z \sigma_n(\lambda) n(z) dz$$

where  $\sigma_n(\lambda)$  is the *cross section* for the absorption process we are looking at,  $n(z)$  is the density of absorbing molecules per volume in the pathway. Note that this is wavelength dependent!

Example: *absorption cross sections* of ozone

## Ex.: Absorption cross section, ozone



## Scattering, Rayleigh

*Rayleigh scattering* is the process of scattering on air molecules. The scattering cross section of air molecules is given by:

$$\sigma_{Ray}(\lambda) = \frac{8\pi}{3} \left( \frac{2\pi}{\lambda} \right)^4 \alpha_p^2$$

What is important here is that this cross section *increase with decreasing wavelength*.

The scattering cross section of air is larger for blue and ultraviolet than for visible radiation.

## Extinction = absorption + scattering

When light is absorbed by a medium of optical thickness the total optical thickness is a sum of all attenuation processes. For absorption and scattering processes:

$$\text{in general: } \tau_t = \tau_{abs} + \tau_{scat}$$

In the atmosphere there are several absorption and scattering processes:

Ozone absorption (uv).

Scattering by clouds, by air and by aerosols.

The spectrum  $F(\lambda, z)$  is modified by the absorption of all atmospheric constituents integrated over the line-of-sight from the satellite to the star, according to the Beer-Lambert law. When only ozone absorption is considered (for simplicity) the spectrum  $m$  can be written as:

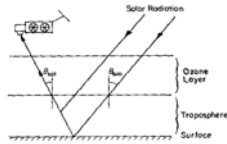
$$F(\lambda, z) = F_o(\lambda) e^{-\sigma_\lambda N(z)}$$

$\lambda$  is the wavelength,  $z$  is the altitude of the line-of-sight above the horizon,  $N(z)$  (molecules/cm<sup>2</sup>) is the integrated quantity of ozone along the line-of-sight and  $\sigma$  the absorption cross section of ozone.

$$N(z) = -\frac{1}{\sigma_\lambda} \log\left(\frac{F(\lambda, z)}{F_o(\lambda)}\right)$$

## Bestemmelse av mengde ozon.

- $F(\lambda) = F_{\text{sun}}(\lambda) \cdot \text{radiance measured by the sat}$
- $F_{\text{sun}}(\lambda) =$  solar irradiance at the top of the atmosphere
- combined reflectance
- $\rho_{\text{Ozone}}$  must be solved
- $\rho_{\text{Ozone}} = \exp[-\sigma_a(\lambda) U_{\text{Ozone}}]$
- $U_{\text{Ozone}}$  (mass of ozone per unit area) = density  $N_{\text{Ozone}}$  • pathlength.
- $N \propto \log_{10}[F(\lambda_1)/F_{\text{sun}}(\lambda_1)]$
- Select two wavelengths  $\lambda_1$  and  $\lambda_2$  close to each other to remove  $\rho_{\text{rest}}$



## Ozone measurements

If we now combine measurements from the two pairs 1 and 2:

$$\left(\frac{I}{P}\right)_1 \left(\frac{I}{P}\right)_2 = \left(\frac{I_0}{P_0}\right)_1 \left(\frac{I_0}{P_0}\right)_2 \left(\frac{\exp(-\tau_1 m)}{\exp(-\tau'_1 m)}\right)_1 \left(\frac{\exp(-\tau_2 m)}{\exp(-\tau'_2 m)}\right)_2$$

If we rearrange and solve this for X as we did just for one pair we get

$$X_{12} = \frac{N_1 - N_2 - [(\tau_R - \tau'_R)_1 - (\tau_R - \tau'_R)_2] m_R \frac{p}{p_0} - [(\tau_a - \tau'_a)_1 - (\tau_a - \tau'_a)_2] m_a}{\mu[(\alpha_{O3} - \alpha'_{O3})_1 - (\alpha_{O3} - \alpha'_{O3})_2]}$$

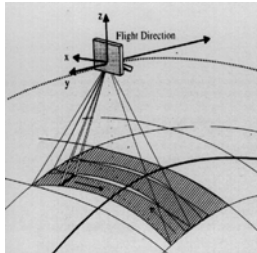
and now if we chose the second in a right manner we can remove the aerosol dependence  $[(\tau_a - \tau'_a)_1 - (\tau_a - \tau'_a)_2] m_a$  and say that this cancel out. This leaves us with

$$X_{12} = \frac{N_1 - N_2 - [(\tau_R - \tau'_R)_1 - (\tau_R - \tau'_R)_2] m_R \frac{p}{p_0}}{\mu[(\alpha_{O3} - \alpha'_{O3})_1 - (\alpha_{O3} - \alpha'_{O3})_2]}$$

Since the air mass of rayleigh scattering and the air mass for rayleigh scattering is well known we only have to concern with the ozone absorption coefficients and ozone air mass  $\mu$ . But these values has been calculated and exist in tables and hence we can get an estimate of the total ozone  $X_{12}$  in the atmosphere above the dobson instrument.

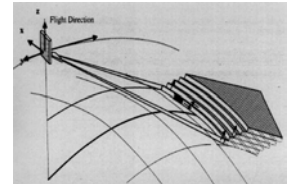
## Ulike måleinstillinger - Nadir

- PRINCIPLE
- The instrument is scanning across-track with a selectable swath width of up to +/- 500 km with respect to the subsatellite track. This maximum swath ensure a global coverage with a maximum of three days at the equator.
- APPLICATIONS
- Determination of total column amounts of trace gazes
- Stratospheric and tropospheric profiles of certain species at limited height resolution
- Cloud detection
- Aerosols
- Surface reflectance



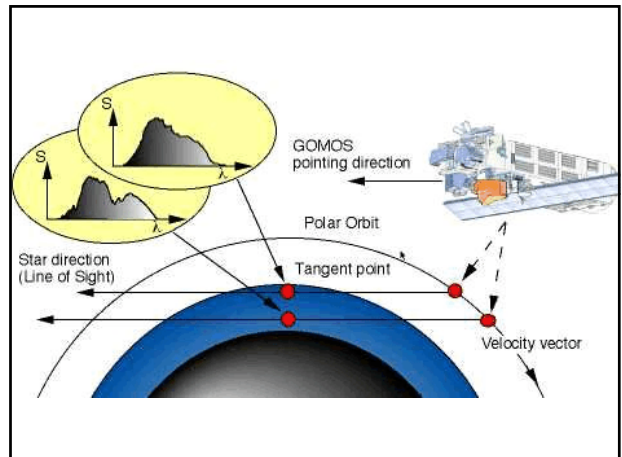
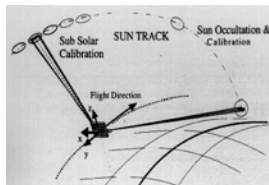
## Ulike måleinstillinger - Limb

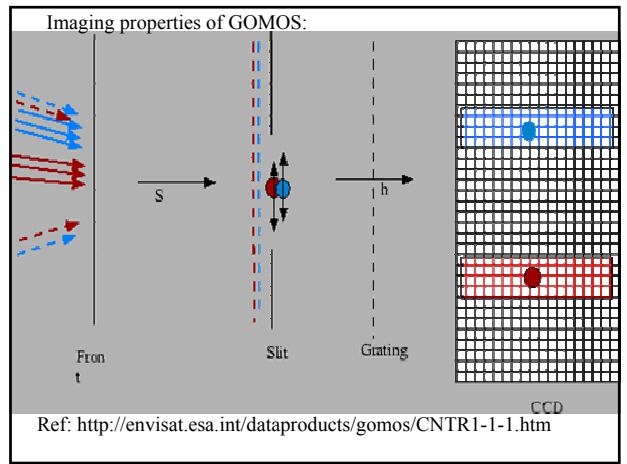
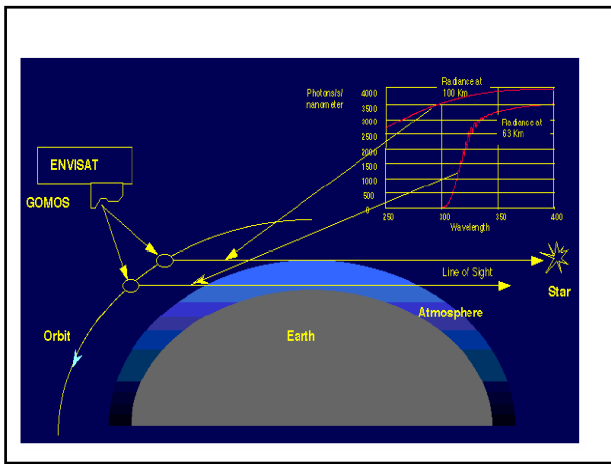
- PRINCIPLE
- Looking forward and starting at the Earth horizon the atmosphere is scanned tangentially over a 1000km swath.
- At the end of each azimuth scan, the elevation is increased by a step which correspond of a increase in altitude at the tangent point of 3 km. This basic scan pattern is reproduced until the maximum altitude of 100 km is reached.
- APPLICATIONS
- Determination of trace gas stratospheric profiles with a improved resolution (compared to nadir)
- Aerosol profiles



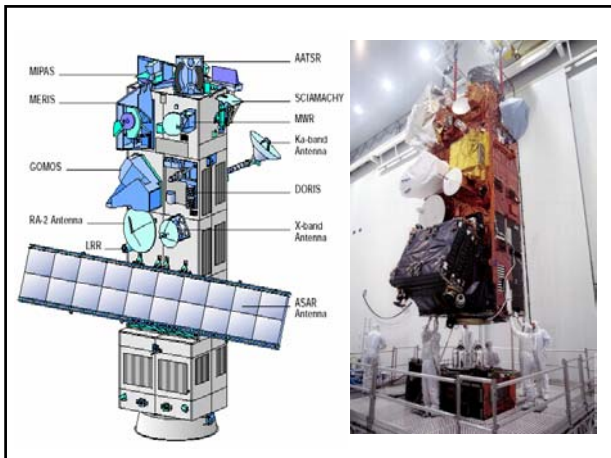
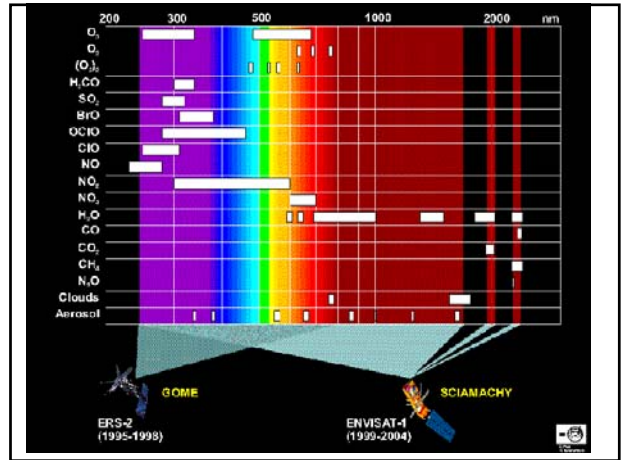
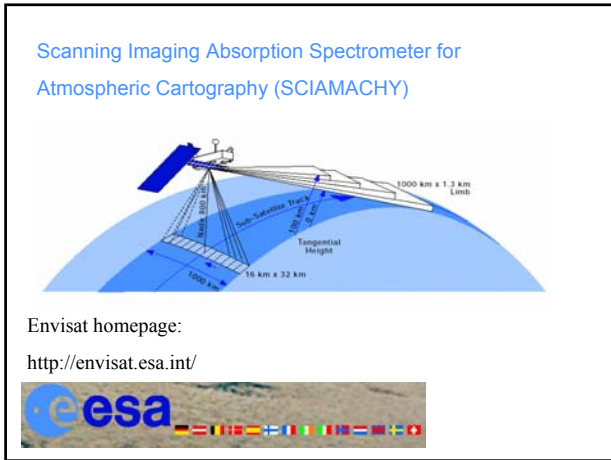
## Ulike måleinstillinger - Occultation

- Special for Gomos, on the new ENVISAT
- Selects a target, the sun or the moon and tracks it as soon as it comes above the horizon and until the line of sight reaches a maximum tangent height of 100 km.
- Spectra of the source radiation transmitted through the atmosphere are recorded during the complete sequence.
- APPLICATIONS
- Possible uses of occultations measurements are quite similar to limb measurements i.e. there are well suited to determine distribution profiles of trace gazes and aerosols.



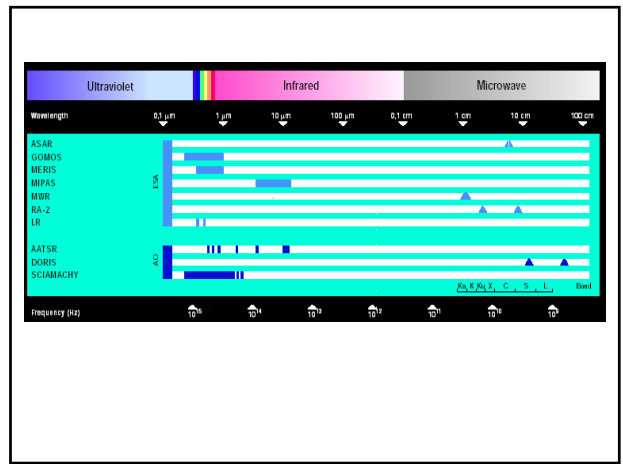
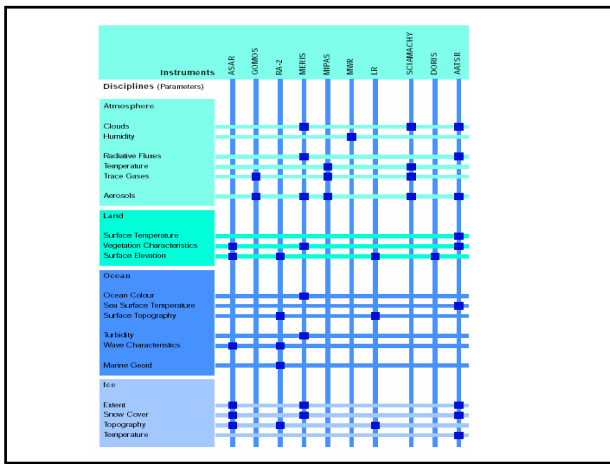


Ref: <http://envisat.esa.int/dataproducts/gomos/CNTR1-1-1.htm>



- Platform:
- ESA Developed Instruments (EDI's)
  - MERIS** (Medium Resolution Imaging Spectrometer)
  - MIPAS** (Michelson Interferometric Passive Atmospheric Sounder)
  - ASAR** (Advanced Synthetic Aperture Radar)
  - GOMOS** (Global Ozone Monitoring by Occultation of Stars)
  - RA2** (Radar Altimeter 2)
  - MWR** (Microwave Radiometer)
  - LRR** (Laser Retro Reflector)





## Observasjon av havet, produktspekteret til MERIS

ESA's nye instrument på ENVISAT.  
<http://envisat.estec.esa.nl/index.html>

- MERIS is designed to acquire 15 spectral bands in the 390 - 1040 nm.
- 15 spectral bands has been derived for oceanographic and interdisciplinary applications. Mulige anvendelser er angitt.

Band centre (nm)	Bandwidth (nm)	Potential Applications
1 412.5	10	Yellow substance, turbidity
2 442.5	10	Chlorophyll absorption maximum
3 490	10	Chlorophyll, other pigments
4 510	10	Turbidity, suspended sediment, red tides

Band centre (nm)	Bandwidth (nm)	Potential Applications
5 560	10	Chlorophyll reference, suspended sediment
6 620	10	Suspended sediment
7 665	10	Chlorophyll absorption
8 681.25	7.5	Chlorophyll fluorescence
9 705	10	Atmospheric correction, red edge
10 753.75	7.5	Oxygen absorption reference
11 760	2.5	Oxygen absorption R-branch
12 775	15	Aerosols, vegetation
13 865	20	Aerosols corrections over ocean
14 890	10	Water vapour absorption reference
15 900	10	Water vapour absorption, vegetation

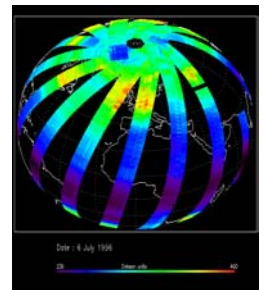
## Produksjon av instrumentering

- Høykvalitets instrumenter, samme prinsipp som brukt ellers.
- Krever utprøving, kalibrering og holdbarhetstester før oppsending
- Kalibreringsrutiner underveis
- Validering via bakkemålinger og fly, ballong ekspedisjoner
- Sammenligning av ulike satellitter



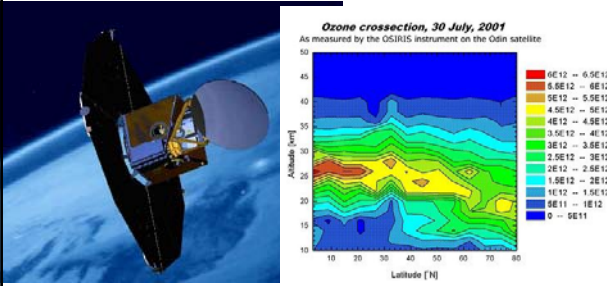
## Different levels of dataproducts

- Raw data: level 1 a
- Data retrieved from calibrated spectral irradiance measurements: level 1b.
- "Non" quality controlled daily images: level 2
- Scientific quality control: level 3 data,
- (usually not official on web).



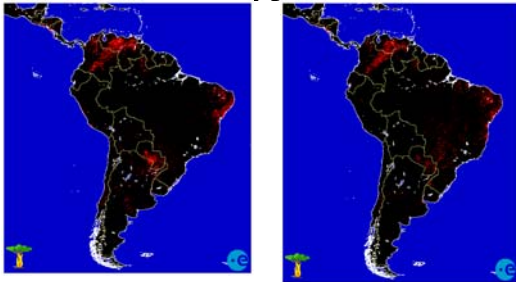
Svensk miljösatellit: Odin, [skutt opp mar](#): 2001

Website: <http://www.ssc.se/ssd/frameodinresults.html>

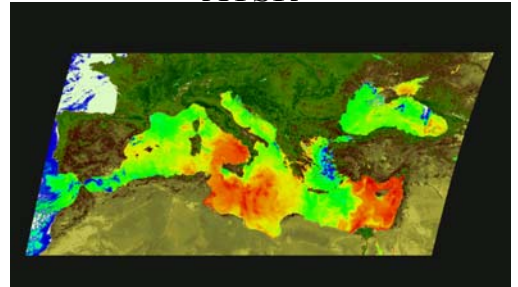


## Jordobservasjon - miljøfysikk

### Fire detection - 1993 January - February ERS-2 ATSR

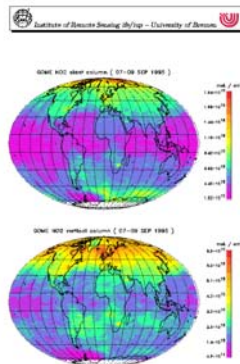


### Sea surface temperatures - ERS-2 ATSR



### ERS: dataproducs

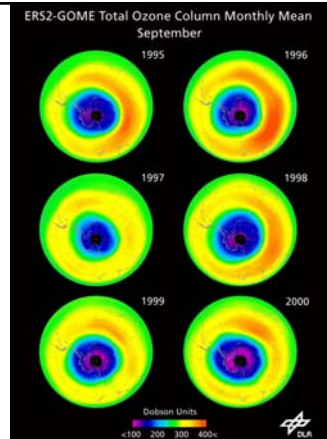
- Total column Ozone
- Ozone profiles
- Nitrogen dioxide, NO<sub>2</sub>
- BrO, OClO, SO<sub>2</sub>
- aerosols (small particles >0.5 μm)
- clouds
- albedo (reflection from ground and clouds)
- photochemistry studies



ERS2 GOME  
dataproducs

September Mean O<sub>3</sub>  
column for southern  
Hemisphere

Product delivered by  
DLR Germany





## SeaWiifs dekning

[seawifs\\_daily\\_coverage\[1\].mpeg](#)



Interessante lenker:

Satelitter

<http://envisat.esa.int> (envisat, homepage)

<http://copi.esa.int> (ESA earth observation , principle investigator portal)

<http://www.ssc.se/ssd> (odin, miljøovervåkning)

<http://eos-chem.gsfc.nasa.gov/> (troposfære kjemi)

<http://jwocky.gsfc.nasa.gov/TOMSmain.html> (ozon)

<http://seawifs.gsfc.nasa.gov/SEAWIFS/SEASTAR/SPACECRAFT.htm>

(hav målinger)

<http://neonet.knmi.nl/> (ozon)

## Litteratur

Introduction to Remote Sensing, 3<sup>rd</sup> Edition, James B. Campbell, The Guilford press, New York London, 2002, 620 pages.